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Full length article

# The characteristics of particulate absorption, scattering and attenuation coefficients in the surface ocean; Contribution of the Tara Oceans expedition



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### ABSTRACT

A dataset consisting of AC-S measurements of (hyper-) spectral particulate absorption, scattering and attenuation coefficients were obtained from measurements performed on the flow-through system of the R/V Tara during its 2.5-year long expedition.

The AC-S instruments were robust, working continuously with weekly maintenance for about 3 months at a time, and provided absorption (attenuation) data for 454 (375) days, or 90% (75%) of total possible days during the expedition.

This dataset has been mapped to 1 km × 1 km bins to avoid over emphasizing redundant data, and to match the spatial scale of typical ocean color satellite sensors. It consists of nearly 70,000 particulate absorption spectra and about 60,000 particulate scattering and attenuation spectra. These data are found to be consistent with chlorophyll extraction and with the published average shapes of particulate absorption and scattering spectra and bio-optical relationships. This dataset is richer than previous ones

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in the data from open-ocean (oligotrophic) environments making it more representative of global distributions and of utility for global algorithm development.

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## 1. Introduction

The Tara Oceans expedition, a 2.5-year long and 57,000 mile long course (Fig. 1), was conceived to provide a snapshot of the distribution of planktonic organisms in the world ocean, providing “A global-scale study of morphological, genetic, and functional biodiversity of plankton organisms in relation to the changing physico-chemical parameters of the oceans” (Karsenti et al., 2011).

As part of this expedition, we installed an AC-S hyper-spectral spectrophotometer (WET Labs, Inc., Philomath, OR) on R/V Tara’s flow-through seawater system, measuring both absorption and attenuation coefficients. The flow-through seawater was supplied by a diaphragm pump plumbed to the bow intake. By alternating sampling with water passing through a 0.2- $\mu\text{m}$  membrane filter cartridge, we obtained the particulate attenuation and absorption coefficients. While this method has been evaluated previously (e.g., Dall’Olmo et al., 2009; Slade et al., 2010; Westberry et al., 2010), it has never been used to this extent.

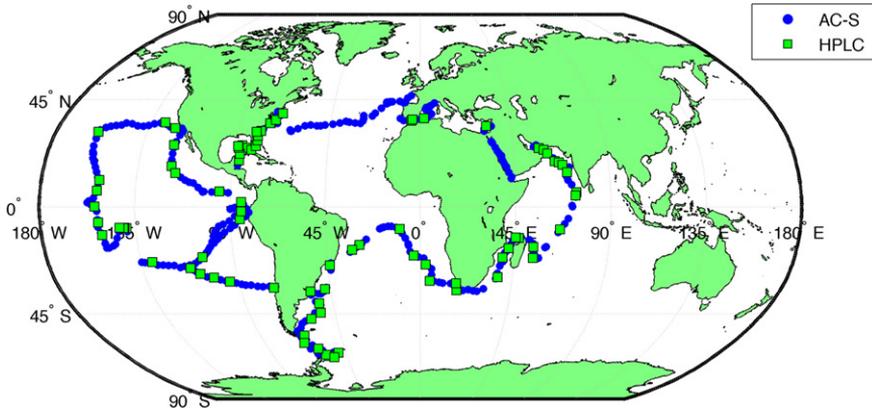
In this manuscript we analyze the robustness of the AC-S, present the salient features of the absorption and attenuation data set obtained, and highlight some of the biogeochemical parameters that can be derived from it. In particular, we compare our hyper-spectral data with the multi-spectral data (measured with WET Labs AC-9 instruments) presented in Barnard et al. (1998) and Babin et al. (2003), the most extensive datasets of particulate IOPs published to date.

## 2. Methods

The in-line sampling system (including a photo) and the data processing methodology have been described in detail in Slade et al. (2010). In short, water from the ship’s flow-through system (pumped from  $\sim 2$  m below surface) was passed through a debubbler and then through a WET Labs AC-S. The AC-S sensor was operated constantly (lamps on) except when in port or once per week for weekly instrument cleaning and filter cartridge replacement. Periodic switching of a custom valve passed 0.2  $\mu\text{m}$  filtered seawater through the sensor for 10 min of every hour, or 5 min of every half-hour in coastal waters (the automated valve system has been recently commercialized by Sequoia Scientific, Inc., Bellevue, WA). Interpolating between the readings of filtered water to the time when measurements of the unfiltered water are made (and their subtraction from the measurements without filter) provided the particulate absorption and attenuation values. Absorption data are then corrected for scattering and residual temperature (Slade et al., 2010). The only exception to the protocol in Slade et al. (2010) was for cases when we had absorption data but no attenuation measurements (due to the failure of the c-side of the AC-S ( $\sim 10\%$  of the sampling days)). In that case we processed the data in two ways: (1) assume scattering to be spectrally flat when scatter-correcting the data, as typically done with bench-top spectrophotometers; (2) use the bio-optical relations derived in this paper to obtain a chlorophyll-based attenuation spectra, and use these attenuation spectra to scatter correct the absorption (see below).

In series with the AC-S, a SBE45 MicroTSG (Sea-Bird Electronics, Inc., Bellevue, WA) unit was installed in the flow-through system. Note that due to the subtraction of the dissolved spectra and the residual temperature correction, the SBE45 temperature and salinity measurements were not necessary for the processing of the flow-through AC-S measurements.

Data were first averaged to 1-min time bins (to average high frequency system noise, e.g. due to particles which are observed differently by the a- and c-sides of the instrument). These 1 min data were submitted to databases (SeaBASS and PANGAEA) as merged files (using time as the common



**Fig. 1.** The locations along the Tara Oceans expedition track where AC-S data are available (blue dots) and where HPLC pigment data were collected (squares).

reference) of the time (synced with GPS), position (GPS), temperature, salinity (TSG) and particulate absorption and attenuation (AC-S).

Second, we further spatially median binned the data into  $1 \text{ km} \times 1 \text{ km}$  bins to provide similar scales to those associated with ocean color satellites. Spatial binning avoided over-weighting relatively large datasets collected over nearly two days during Tara Oceans stations conducted near one location while the in-line system was on. The  $1 \text{ km} \times 1 \text{ km}$  global grid was generated using the SeaWiFS equal area gridding scheme (Campbell et al., 2013). In the analysis requiring uncertainties we only use  $1 \text{ km}^2$  bins where at least three 1 min spectra were located, and their standard deviation is used as the measure of uncertainty.

Seawater was collected near the surface (depth  $< 5.1 \text{ m}$ ) for HPLC pigment analysis at 85 locations where we had AC-S measurements along the cruise track of the R/V Tara. For each sample, we vacuum-filtered 2 L of water through 25 mm (in some cases, 47 mm) diameter Whatman GF/F glass filters with  $0.7 \mu\text{m}$  capacity. Filters were frozen in liquid nitrogen and shipped back at the end of up to three legs for extraction and HPLC analysis at the Laboratoire d'Océanographie de Villefranche. The filters were extracted in 100% methanol (3 mL for 25-mm filters, 6 mL for 47-mm filters) for 2 h, disrupted by sonication, and then clarified by filtration (Whatman GF/F). HPLC analysis was carried out within 24 h after extraction, according to the method described by Ras et al. (2008) adapted from Van Heukelem and Thomas (2001). Here we use only the Total Chlorophyll *a* value (= "true" Chlorophyll *a* + Divinyl-Chlorophyll *a* + Chlorophyllide *a*) in order to calibrate the chlorophyll *a* pigment product derived from the AC-S.

To compute averaged spectra for different trophic regions, we divided the dataset into those with chlorophyll *a* (as determined by AC-S absorption, see below) concentration below  $0.1 \text{ mg m}^{-3}$  (oligotrophic), those with chlorophyll *a* concentration between  $0.1$  and  $1 \text{ mg m}^{-3}$  (mesotrophic) and those with chlorophyll *a* concentration greater than  $1 \text{ mg m}^{-3}$  (eutrophic). We then computed a weighted-average spectra using, as weight, the inverse of the standard deviation of all the spectra wavelengths taken within a  $1 \text{ km} \times 1 \text{ km}$  bin, using only spatial bins with at least three 1-min spectra. Finally we also normalized the spectra by area as in Babin et al. (2003), for the ease of spectral-shape comparison.

### 2.1. Derived biochemical parameters

Using the hyperspectral particulate absorption and attenuation spectra three biogeochemical parameters were derived:

1. a particle size index ( $\gamma$ , Boss et al., 2001), by non-linearly fitting a power-law to the particulate beam attenuation spectra ( $\gamma$  is the negative of the power-law exponent). This parameter is negatively

correlated with the mean particle size for small particles less than approximately 20  $\mu\text{m}$ ; the smaller this parameter is the larger the mean particle size. If the size distribution can be assumed to obey a power-law distribution as a function of size (e.g. Junge-type distribution), the exponent of the particle number size distribution ( $\xi$ ) is related to ( $\gamma$ ) by  $\xi = \gamma + 3 - 0.5 \exp(-6\gamma)$  (Boss et al., 2001).

We similarly compute the power-law exponent of particulate scattering ( $\gamma_b$ ). The  $\gamma$ 's are obtained by non-linear least-square fit of the power-law function to the particulate attenuation or scattering spectra using the uncertainty at each wavelength as a weight (thus weighing less wavelengths where the uncertainty in the attenuation or scattering is large). Note that while  $\gamma$  is independent of the scattering correction scheme used to obtain the particulate absorption,  $\gamma_b$  could be affected by it (as scattering is computed as the difference between attenuation and absorption);

2. a total chlorophyll *a* concentration product ( $C_{a\text{-Tot}}$ ) is derived, inspired by Bricaud et al. (1995). We compute the absorption line height at 676 nm as in Boss et al. (2007) to approximate the phytoplankton absorption at this wavelength:

$$a_\phi(676) = [a_p(676) - 39/65a_p(650) - 26/65a_p(715)] (\text{m}^{-1}).$$

For this comparison we averaged the 1-min binned data taken within  $\pm 30$  min of HPLC sampling.  $C_{a\text{-Tot}}$  ranged from 0.011 to 7.27  $\text{mg m}^{-3}$ .

We linearly regressed (type-I)  $\log(a_\phi(676))$  against  $\log(C_{a\text{-Tot}})$  using all the HPLC data within 5.1 m from the surface ( $N = 85, 7.27 \geq C_{a\text{-Tot}} \geq 0.011 \text{ mg m}^{-3}$ , Fig. 2) and obtained the following relationship (with 95% confidence interval in brackets):

$$\log(a_\phi(676)) = 0.82(+/-0.08) \log(C_{a\text{-Tot}}) - 4.14(+/-0.16),$$

$$\text{equivalent to: } a_\phi(676) = 0.016(+/-0.003)C_{a\text{-Tot}}^{0.82+/-0.08} (\text{m}^{-1}),$$

which is inverted to obtain a proxy for chlorophyll *a* from  $a_\phi(676)$ :

$$C_{a, AC-S} = 157a_\phi(676)^{1.22} (\text{mg m}^{-3}).$$

The correlation coefficient between the logarithms of  $C_{a\text{-Tot}}$  and  $a_\phi(676)$  is 0.96 and the RMS relative error in predicting  $C_{a\text{-Tot}}$  from  $a_\phi(676)$  is 38%.

Note that in the cases where no attenuation data were available we obtained this parameter only using the spectrally flat scattering correction, as chlorophyll is a needed input for the alternative one (see below). This is unlikely to result in significant error as the calculation of  $C_{a, AC-S}$  is based on wavelengths where the difference between the scattering correction methods is minimal (it is most significant in blue wavelengths);

3. a particulate organic carbon (POC) estimate (using Global relationship from Gardner et al. (2006)).

$$\text{POC} = 380c_p(660) (\text{mg C m}^{-3}).$$

Given the difference from other published relations (e.g. Cetinić et al. (2012) and references therein, taking into account variations in instruments and wavelength), the uncertainty in POC is likely in the order of  $\pm 50\%$ .

## 2.2. Scattering correction in cases where no attenuation was available

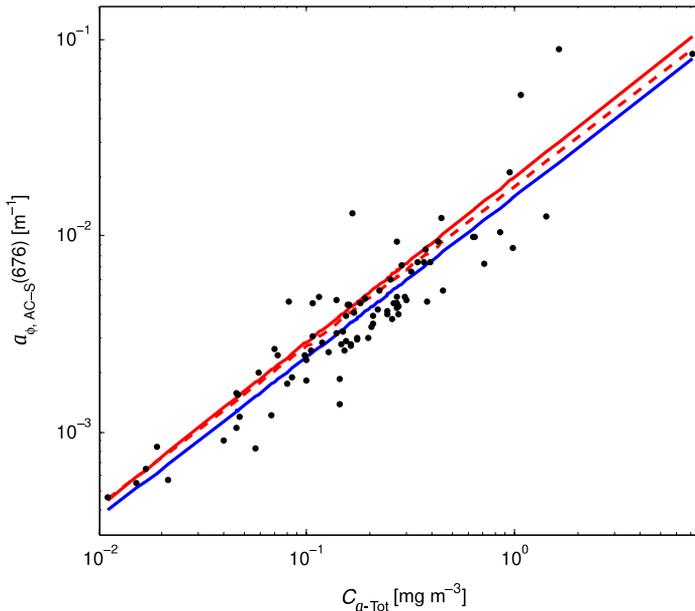
In such cases we applied two scattering corrections: (1) assuming spectrally flat scattering spectra as used with bench-top spectrophotometers; and (2) computing the attenuation spectra based on the bio-optical relations derived here:

$$c_p(488) = 0.377C_{a, AC-S}^{0.53} (\text{m}^{-1}).$$

$$\gamma = 1(C_{a, AC-S} < 0.17) \quad \text{or} \quad \gamma = 0.7 - 0.17 * \log(C_{a, AC-S})(C_{a, AC-S} \geq 0.17)$$

$$c_p(\lambda) = c_p(488)\{\lambda/488\}^{-\gamma}.$$

This spectra was then used, as in Slade et al. (2010), to scatter correct the particulate absorption spectra. In all cases we assumed the particulate absorption spectra to have a zero average for wavelengths between 726–730 nm.



**Fig. 2.** AC-S derived absorption line-height at 676 nm vs. total chlorophyll *a* concentration as determined by HPLC. Blue line is the regression obtained here while the red line is based on Bricaud et al. (1995) and dashed red line is based on Bricaud et al. (1998). The correlation coefficient between the logarithms of  $C_{a-Tot}$  and  $a_{\phi, AC-S}$  is 0.96 and the rms relative error is 38%.  $N = 85$ . (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

**Table 1**

Number of available 1 km<sup>2</sup> measurements of particulate absorption, scattering and attenuation at specific wavelengths and their respective statistics.  $a_{ps}$  denotes the particulate absorption for which we used the chlorophyll based scattering correction when the attenuation was not available (see Section 2).

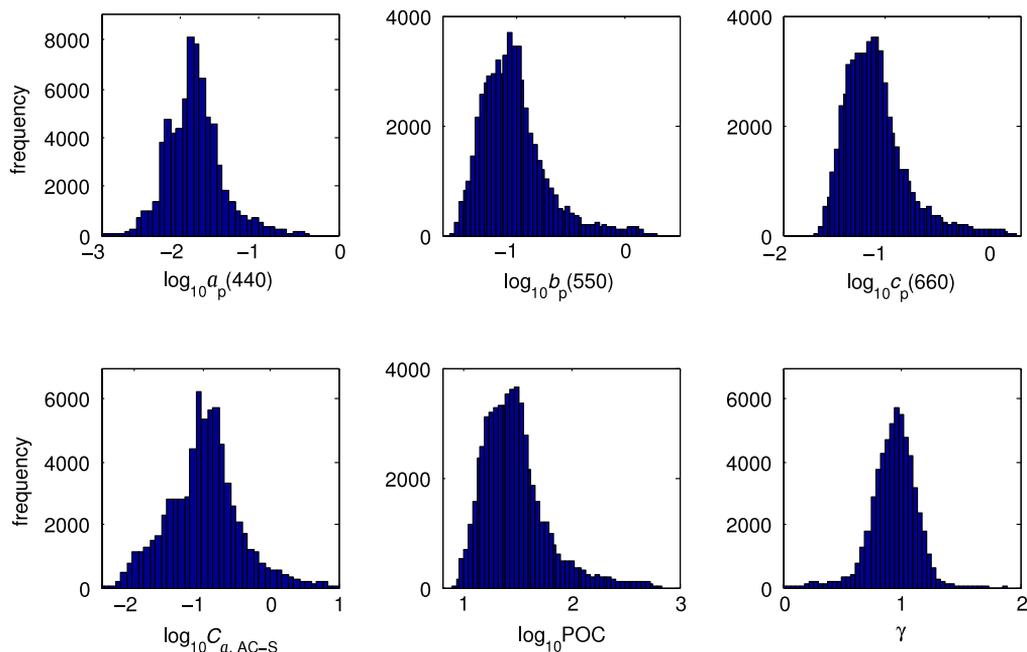
IOP at representative $\lambda$	$N$	Mean	Median	0.5%	2.5%	97.5%	99.5%
$a_p$ (440) ( $m^{-1}$ )	69,944	0.022	0.015	0.003	0.004	0.098	0.216
$a_{ps}$ (440) ( $m^{-1}$ )	69,944	0.022	0.014	0.003	0.003	0.097	0.221
$b_p$ (550) ( $m^{-1}$ )	55,883	0.128	0.085	0.029	0.033	0.609	1.280
$c_p$ (660) ( $m^{-1}$ )	57,188	0.111	0.072	0.025	0.029	0.530	1.196

The resulting difference in  $a_p(440)$  between the two corrections (flat-proportional) varied from 0.001 to 0.009  $m^{-1}$  at the 2.5% and 97.5%, respectively, with a median of 0.004  $m^{-1}$ . The relative difference in  $a_p(440)$  between the two corrections  $\{2|\text{flat-proportional}|/(\text{flat} + \text{proportional})\}$  varied from 1.6% to 67% at the 2.5% and 97.5% level, respectively, with a median of 25%. For the full data set, however, the statistics (Table 1) and statistical relations were not significantly different and we therefore display only the ones with the spectrally flat correction.

### 3. Results and discussion

#### 3.1. Instrumental considerations

The vessel flow-through system was operational a total of 552 days. We lost 51 days of AC-S data in the Mediterranean due to a problem with the data backup system, resulting in 501 potential sampling days. Of these we have AC-S data for 454 days (90%) of which 73 days are only with absorption (failure of the lamp on c-side) and 6 with only attenuation (failure of the lamp on a-side). We had the full set of particulate properties for 375 days (75%). There were two cases where AC-S malfunctions were fixed by the resident engineer on the vessel (in consultation with WETLabs).



**Fig. 3.** Frequency distributions of selected IOPs and parameters derived from the AC-S spectra. Statistics associated with these distributions are in Tables 1 and 2.

We used 5 different AC-S sensors belonging to four different labs (serial numbers S/N 007, 043, 057, 082, 091) operating for nine different periods through the 2.5 years of the expedition (average of 3.3 months per sensor, approximately 3 cruise legs, with typical break of  $\sim 1$  week between legs). After each deployment the AC-S lamps were replaced at WET Labs and the sensor was recalibrated.

Given that the tables for temperature corrections were created for the S/N 007 sensor (Sullivan et al., 2006), it was not evident that they would apply for other sensors. However, the lack of obvious spectral features in the range of 700–750 nm (see below) suggests using those tables for other sensors (or the same sensor after it had been rebuilt) is sensible.

We obtained a little more than 310,000 1 min averaged particulate absorption spectra and nearly 70,000  $1 \text{ km}^2$  averaged spectra. For comparison, the ‘global’ data set in Barnard et al. (1998) had a total over 1000 profiles (unique locations), which were reduced to 89 ‘representative’ vertical profiles that included 1914 depth-binned spectra.

### 3.2. Distributions of values

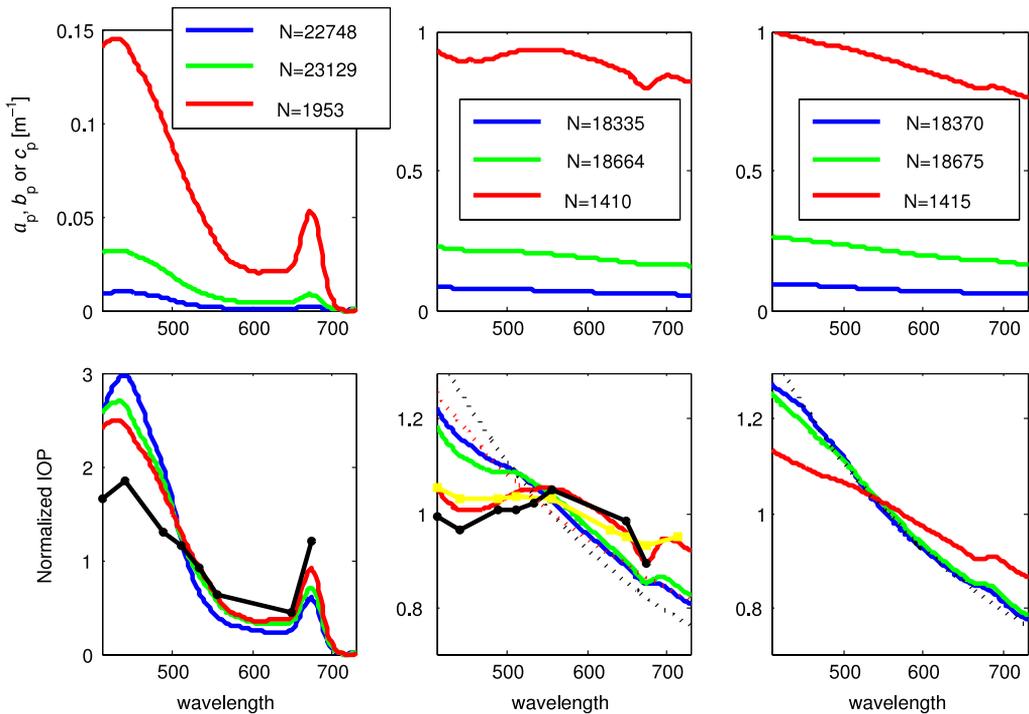
The distributions of IOPs and biogeochemical variables obtained are within the range reported in the literature (Fig. 3, Tables 1 and 2).  $C_{a, AC-S}$  spans three orders of magnitude while POC spans nearly two. The distributions are skewed and do not conform to a parametric log-normal distribution, which is often assumed for bio-optical variables (e.g. Campbell, 1995). This may reflect the fact that these distributions are the sum of different particles (each kind possibly log-normally distributed) of different composition and optical properties.

This data set is particularly rich in oligotrophic data, compared to databases designed for remote sensing applications (Werdell and Bailey, 2005). The value of the spectral exponent of  $c_p$  is relatively constant and has an average spectral shape proportional to  $\lambda^{-1}$  (Fig. 4). Note, however, that the exponent of the power-law fit to particulate scattering (which is the same as that of backscattering if the particulate ratio is constant spectrally, e.g. Whitmire et al. (2007)) is significantly lower (due to the

**Table 2**

Number of available 1 km<sup>2</sup> estimates of total chlorophyll *a*, Particulate organic carbon (POC), the exponent of the power-law fit to the particulate attenuation spectra ( $\gamma$ ), the exponent of the power-law fit to the particulate scattering spectra ( $\gamma_b$ ), and their respective statistics.

Derived parameter	<i>N</i>	Mean	Median	0.5%	2.5%	97.5%	99.5%
Chlorophyll <i>a</i> (mg m <sup>-3</sup> )	69,943	0.26	0.12	0.01	0.013	1.55	4.75
POC (mg m <sup>-3</sup> )	57,188	42.2	27.2	9.5	11.1	201	358
$\gamma$	57,188	0.94	0.95	0.23	0.51	1.25	1.39
$\gamma_b$	55,883	0.73	0.75	-0.08	0.23	1.03	1.18



**Fig. 4.** Averaged particulate absorption (left), scattering (middle) and attenuation (right) spectra for  $C_{a,AC-S} < 0.1$  (blue),  $0.1 < C_{a,AC-S} < 1$  (green) and  $1 < C_{a,AC-S}$  (red) (top panel) and their normalized shapes (bottom panels). Black solid lines are the normalized spectra for the “global” dataset of Barnard et al. (1998), and the yellow line is from Babin et al. (2003). Black dashed line is the normalized  $\lambda^{-1}$  curve while the red dashed line is the normalized  $\lambda^{-0.75}$  curve, where  $\lambda$  denotes the wavelength in vacuum. *N* represents the number of spectra that were uncertainty-weight-averaged to provide the one displayed. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

effect of absorption). In addition, the beam attenuation spectrum seems to deviate from the power-law spectrum near the 676 nm absorption peak and at wavelengths of highly varied absorption near 440 nm. Such a response to absorption has not been resolved previously with AC-9 measurements, and is consistent with anomalous dispersion (van de Hulst, 1957).

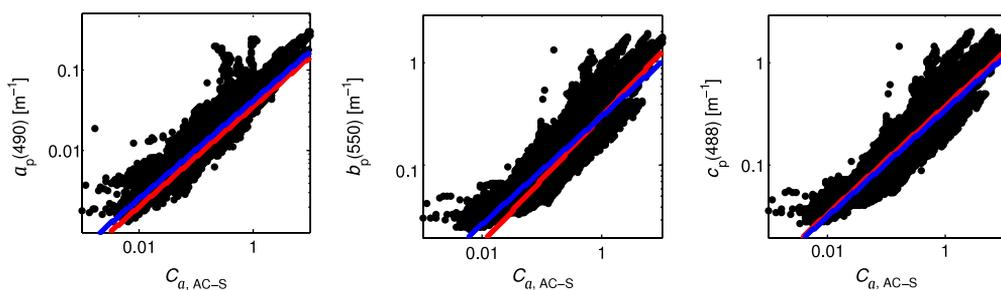
### 3.3. Comparison to published bio-optical relationships and databases

The derived relationship between absorption line height and chlorophyll ( $a_\phi(676) = 0.016 (+/-0.003) C_{a-Tot}^{0.82 +/- 0.08}$ , see Section 2) is similar to Bricaud et al. (1995), who found  $a_\phi(676) = 0.02 C_{a-Tot}^{0.841}$ , and Bricaud et al. (1998) who found  $a_\phi(676) = 0.018 C_{a-Tot}^{0.816}$ , obtained with a bench-top

**Table 3**

Number of available 1 km<sup>2</sup> estimates of measurements of particulate absorption, scattering and attenuation at specific wavelengths used to derive a relationship with AC-S estimated total chlorophyll *a* compared with the relationship provided in Morel (2009). Correlation coefficient ( $\rho$ ) and root-mean-square relative error apply to the equation derived for the Tara dataset. 95% confidence intervals of regression parameters are smaller or equal to the uncertainties implied in the rounding of these parameters. Note that using the two different scattering correction for the cases where attenuation was not available (see text) resulted in regressions for  $a_p$  (490) vs. chlorophyll that were not significantly different.

IOP regressed with $C_{a, AC-S}$	<i>N</i>	Equation here	Morel (2009)	$\rho$	RMSE%
$a_p$ (490) (m <sup>-1</sup> )	69,943	$0.041 C_{a, AC-S}^{0.61}$	$0.034 C_{a-Tot}^{0.62}$	0.97	27
$b_p$ (550) (m <sup>-1</sup> )	55,883	$0.30 C_{a, AC-S}^{0.53}$	$0.30 C_{a-Tot}^{0.62}$	0.95	30
$c_p$ (488) (m <sup>-1</sup> )	55,883	$0.345 C_{a, AC-S}^{0.52}$	$0.377 C_{a-Tot}^{0.53}$	0.95	28



**Fig. 5.** Particulate absorption, scattering and attenuation at selected wavelengths compared to AC-S derived Total Chlorophyll *a*. Relationship derived in this paper (blue) while in red the relationships provided in Morel (2009). Fit equations and their statistics are in Table 3. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

spectrophotometer (Fig. 2), as well as Claustre et al. (2000) who found  $a_p(676) = 0.019C_{a-Tot}^{0.806}$  using an AC-9, providing an additional validation for the AC-S line-height method to obtain  $C_{a-Tot}$  (Fig. 2).

Following Morel (2009) and references therein we regressed the logarithms of particulate absorption, scattering and attenuation at selected wavelengths vs.  $C_{a, AC-S}$  (Table 3). Relationships are similar to those in Morel (2009) for all parameters (Fig. 5). We note that at very low  $C_{a, AC-S}$  the scattering and attenuation coefficient seem to reach a relatively constant value. Such a behavior was observed in Ocean Color inversion of particulate backscattering and may represent the optical properties of “background” non-algal particles (Behrenfeld et al., 2005).

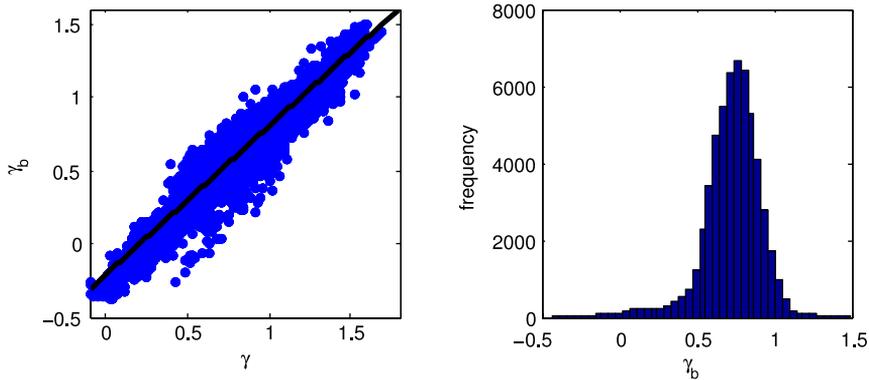
### 3.4. Relation between particulate attenuation and scattering power-law fit exponents

The spectrum of particulate attenuation and scattering do not have the same goodness of fit to a power-law function and their spectral slope is different (Fig. 4). However, recent algorithms for Ocean Color inversion have attempted to obtain the power-law exponent of the particulate backscattering coefficient (Loisel et al., 2006; Kostadinov et al., 2009), which should be the same exponent as that for particulate scattering, if the particulate backscattering ratio is spectrally independent (an assumption that is consistent with measurements, e.g. Whitmire et al. (2007)).

Comparing the power-law exponents for particulate attenuation and scattering (Fig. 6) we found both to be highly correlated ( $r = 0.95$ ) and with a type-II regression relationship (using E. Peltzer’s sqfitma.m, <http://www.mbari.org/staff/etp3/regress/index.htm>) of:

$$\gamma = 0.9886(+/-0.002)\gamma_b + 0.215(+/-0.002)$$

where the terms in parentheses are twice the standard deviation (e.g. 95% confidence interval). The RMSE of this regression is 0.06. In addition, the range of  $\gamma$  and  $\gamma_b$  observed (Table 2) is significantly more restricted than that obtained from remote sensing inversions (e.g. Fig. 4 in Loisel et al. (2006) and Fig. 6 in Kostadinov et al. (2009)).



**Fig. 6.** Exponent of the power-law fit to the particulate scattering spectra ( $\gamma_b$ ) as a function of the exponent of the power-law fit to the particulate attenuation spectra ( $\gamma$ , left) and the frequency distribution of  $\gamma_b$  (right). Black line denotes the type II least-squares fit (details in text).

#### 4. Conclusions

The data collected on board the R/V Tara demonstrate that it is possible, using minimal maintenance (weekly cleaning and filter replacement), to obtain high quality particulate absorption and attenuation at resolutions  $< 0.001 \text{ m}^{-1}$ .

The AC-S provided highly resolved optical properties and biogeochemical proxies throughout the expedition, on time scales of several months, the time scale of a typical oceanic expedition. Our results (flat absorption spectra for  $\lambda > 700$  following residual temperature correction) support the use of Sullivan et al. correction tables (Sullivan et al., 2006) for all AC-Ss to account for the effects of temperature (though more testing is needed as, given our method, only small temperature corrections needed to be applied).

The Tara expedition has collected the most ‘global’ and self-consistent set of particulate absorption, scattering and attenuation spectra to date consisting of nearly 70,000 (60,000) spectra of absorption (attenuation and scattering). This dataset differs significantly from previous ones (Barnard et al., 1998; Babin et al., 2003) in that it contains significantly more data in oligotrophic waters (e.g. Fig. 4) and agrees well with the more globally representative dataset used in the previous Case 1 studies (Morel, 2009). As with the older datasets, we expect this new dataset to be useful for parameterization of IOPs in radiative transfer modeling of the ocean. In addition, it provides a wide range of conditions to be used for the validation of remote sensing algorithms (see companion paper by Werdell et al. (2013)). Frequency distributions of particulate IOPs and derived variables are not consistent with a log-normal distribution, in contrast to often-used parameterizations (e.g. Campbell (1995) but see Whitmire et al. (2007)).

The data collected are consistent with the previous observations for waters with similar values of chlorophyll (Barnard et al., 1998; Babin et al., 2003) as well as for waters with a wide range of chlorophyll concentrations (Bricaud et al., 1995, 1998; Claustre et al., 2000; Morel, 2009). The spectrum of beam attenuation, while well represented over most of the samples by a power-law spectrum with an exponent near  $-1$  is, none-the-less, perturbed from the power-law shape where absorption peaks are significant (anomalous dispersion).

The scattering spectra are not well represented by such a shape over a wider range of wavelength (which is still the most prevalent model, e.g. IOCCG (2006)). However, when fitting such a power-law shape for particulate scattering, its value is highly correlated with that of the spectral particulate attenuation and its exponent is offset by  $-0.21$  (RMSE of  $\pm 0.06$ ). This suggests that the particulate backscattering spectra, if retrievable from the Ocean Color, could, indeed, provide information about size tendencies (yet validation of this concept on oceanic scales is lacking).

The consistency between the bio-optical relationship within the Tara dataset and those obtained previously (Morel, 2009) using widely different instrumentation and processing, suggests that the

Tara dataset can be used for remote-sensing validation activities. Indeed, Werdell et al. (2013) find them to be consistent with the previous validation data.

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